

# Photodesorption of $\text{H}_2\text{O}$ , $\text{HDO}$ , and $\text{D}_2\text{O}$ ice and its impact on fractionation $\star, \star\star$ (Research Note)

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## ABSTRACT

The  $\text{HDO}/\text{H}_2\text{O}$  ratio measured in interstellar gas is often used to draw conclusions on the formation and evolution of water in star-forming regions and, by comparison with cometary data, on the origin of water on Earth. In cold cores and in the outer regions of protoplanetary disks, an important source of gas-phase water comes from photodesorption of water ice. This research note presents fitting formulae for implementation in astrochemical models using previously computed photodesorption efficiencies for all water ice isotopologues obtained with classical molecular dynamics simulations. The results are used to investigate to what extent the gas-phase  $\text{HDO}/\text{H}_2\text{O}$  ratio reflects that present in the ice or whether fractionation can occur during the photodesorption process. Probabilities for the top four monolayers are presented for photodesorption of  $\text{X}$  ( $\text{X}=\text{H}, \text{D}$ ) atoms,  $\text{OX}$  radicals, and  $\text{X}_2\text{O}$  and  $\text{HDO}$  molecules following photodissociation of  $\text{H}_2\text{O}$ ,  $\text{D}_2\text{O}$ , and  $\text{HDO}$  in  $\text{H}_2\text{O}$  amorphous ice at ice temperatures from 10–100 K. Significant isotope effects are found for all possible products: (1) H atom photodesorption probabilities from  $\text{H}_2\text{O}$  ice are larger than those for D atom photodesorption from  $\text{D}_2\text{O}$  ice by a factor of 1.1; the ratio of H and D photodesorbed upon  $\text{HDO}$  photodissociation is a factor of 2. This process will enrich the ice in deuterium atoms over time; (2) the  $\text{OD}/\text{OH}$  photodesorption ratio upon  $\text{D}_2\text{O}$  and  $\text{H}_2\text{O}$  photodissociation is on average a factor of 2, but the  $\text{OD}/\text{OH}$  photodesorption ratio upon  $\text{HDO}$  photodissociation is almost constant at unity for all ice temperatures; (3) D atoms are more effective in kicking out neighbouring water molecules than H atoms. However, the ratio of the photodesorbed  $\text{HDO}$  and  $\text{H}_2\text{O}$  molecules is equal to the  $\text{HDO}/\text{H}_2\text{O}$  ratio in the ice, therefore, there is no isotope fractionation when  $\text{HDO}$  and  $\text{H}_2\text{O}$  photodesorb from the ice. Nevertheless, the enrichment of the ice in D atoms due to photodesorption can over time lead to an enhanced  $\text{HDO}/\text{H}_2\text{O}$  ratio in the ice, and, when photodesorbed, also in the gas. The extent to which the ortho/para ratio of  $\text{H}_2\text{O}$  can be modified by the photodesorption process is discussed briefly as well.

**Key words.** astrochemistry – molecular processes – ISM: molecules – solid state: volatile

## 1. Introduction

The abundance of water in interstellar clouds, its partitioning between gas and ice, and its evolution from collapsing cores to protoplanetary disks are topics that are directly relevant for the formation of planets and their atmospheres. Astronomers can study the water trail starting either from the very first stages of dense molecular clouds even before a star is born (Whittet et al., 1988; Boogert et al., 2008; Caselli et al., 2012) or working backwards from the oceans on our own planet Earth, trying to link its composition to that of icy solar system bodies (e.g., Nuth 2008, Mumma & Charnley 2011, and see summary in van Dishoeck et al. 2014). In both approaches, the  $\text{HDO}/\text{H}_2\text{O}$  ratio is thought to be a good indicator of the relevant processes involving water. Specifically, the fact that the  $\text{HDO}/\text{H}_2\text{O}$  abundance of  $3.1 \times 10^{-4}$  in Earth's oceans (de Laeter et al., 2003) is a factor of 2 lower than that found in most comets (Villanueva et al., 2009) has been

used as an argument that most water on Earth originates from impact of asteroids rather than comets (Raymond et al., 2004)<sup>1</sup>. Recent observations with the *Herschel Space Observatory* have found at least two comets for which the  $\text{HDO}/\text{H}_2\text{O}$  ratio is nearly the same as that in Earth's oceans, demonstrating that there is a population of icy bodies with similar values (Hartogh et al., 2011; Bockelée-Morvan et al., 2012). However, 'hot off the press' data from the *Rosetta* mission to comet 67 P/Churyumov-Gerasimenko now suggest that the  $\text{HDO}/\text{H}_2\text{O}$  in Jupiter family objects has a wider range than heretofore assumed triggering further debate on the origin of Earth's water (Altwegg et al., 2014). Hence, there is increased interest in understanding what processes affect the  $\text{HDO}/\text{H}_2\text{O}$  ratio from core to disk to icy planetary bodies.

Physical-chemical models of the collapse of interstellar clouds and disk formation suggest that the bulk of water is formed on grains in the dense pre-collapse stage and enters the disk as ice that is largely unaltered since its formation (e.g., Aikawa & Herbst, 1999; Visser et al., 2009). If so, the  $\text{HDO}/\text{H}_2\text{O}$  ratio in the disk (and eventually in the planetesimals and planets

<sup>\*</sup> Appendices are available in electronic form at <http://www.aanda.org>

<sup>\*\*</sup> Compiled simulation data and raw data are only available at the CDS via anonymous ftp to [cdsarc.u-strasbg.fr](http://cdsarc.u-strasbg.fr) (130.79.128.5) or via <http://cdsarc.u-strasbg.fr/viz-bin/qcat?J/A+A/> vol/page

<sup>1</sup> The measured  $\text{HDO}/\text{H}_2\text{O}$  ratio is  $2 \times (\text{D}/\text{H})$  in water; the latter values are commonly used in the literature.

that formed from it) should reflect the ice values found in the interstellar cloud from which the star formed. Recent simulations of deuterium chemistry in protoplanetary disks show that, in the absence of strong vertical mixing, a high initial HDO/H<sub>2</sub>O ratio is required in the ice to reproduce the values seen in solar system bodies (Furuya et al., 2013; Albertsson et al., 2014; Cleaves et al., 2014). In cold clouds prior to star formation and in the cold outer envelopes of low-mass protostars ( $T < 20$  K), high ratios of deuterated species up to a few  $\times 10^{-2}$  have been found for many molecules including water (e.g., van Dishoeck et al. 1995, Ceccarelli et al. 1998, Bacmann et al. 2003, Parise et al. 2004, Parise et al. 2012 and see review by Ceccarelli et al. 2014). Although the fractionation of HDO is generally not as high as that of other molecules, values of HDO/H<sub>2</sub>O up to a few % have been inferred in cold gas (Liu et al., 2011; Coutens et al., 2012), higher than those found in comets. They are also high compared with the upper limits on HDO/H<sub>2</sub>O measured directly in interstellar ices of  $< (1-5) \times 10^{-3}$  (Dartois et al., 2003; Parise et al., 2003). Moreover, the HDO/H<sub>2</sub>O ratio in warm gas ( $> 100$  K) near protostars typically has lower values of  $\sim 10^{-3}$ , within a factor of a few of those of comets (e.g., Gensheimer et al., 1996; van der Tak et al., 2006; Jørgensen & van Dishoeck, 2010b; Persson et al., 2014; Emprechtinger et al., 2013; Neill et al., 2013; Coutens et al., 2014a,b). These differences raise the question whether the high HDO/H<sub>2</sub>O values measured in cold gas reflect the HDO/H<sub>2</sub>O value in the ice or whether they have been altered by desorption processes.

In cold clouds, the bulk of the water is present as ice with only a small fraction of water in the gas, typically  $10^{-4}$  of that of ice (e.g., Boonman & van Dishoeck, 2003; Caselli et al., 2012). The water ice is formed in an early stage of cloud formation once the density and extinction reach a threshold (Cuppen & Herbst, 2007). Temperatures in these regions are too low for thermal sublimation of ice, so the observed cold gas-phase water must come from non-thermal desorption processes of water ice with photodesorption a leading candidate (Hollenbach et al., 2009; Öberg et al., 2009). Indeed, both in pre-stellar cores (Caselli et al., 2012), in the outer parts of protostellar envelopes (Mottram et al., 2013; Schmalzl et al., 2014), and in the cold outer parts of disks (Dominik et al., 2005; Hogerheijde et al., 2011), the small observed gas-phase water amount is consistent with being due just to the photodesorption of water ice. The required UV photons are provided by the general interstellar radiation field at the outside of the cloud or by the star in the case of a protoplanetary disk, and by the interaction of cosmic rays with H<sub>2</sub> deeper inside the cloud. Other non-thermal processes such as chemical desorption are poorly quantified but unlikely to be significant for thick layers of water ice (Dulieu et al., 2013; Minissale & Dulieu, 2014). Chemical models have so far considered mostly deuterium fractionation processes in the gas phase or on the grains, but not during the (photo)desorption process (e.g., Tielens 1983, Aikawa & Herbst 1999, Roberts et al. 2003, Cazaux et al. 2011, Du et al. 2012, Aikawa et al. 2012, Taquet et al. 2013 and see summary in van Dishoeck et al. 2013).

In this paper, we use previously computed photodesorption probabilities to investigate the extent to which the photodesorption process can modify the gas-phase HDO/H<sub>2</sub>O ratio compared with that of the ices. Experiments have demonstrated that the water photodesorption yield is typically  $10^{-3}$  per incident photon, but those data did not have the accuracy to establish differences between H<sub>2</sub>O and its isotopologues (Westley et al., 1995; Öberg et al., 2009). Photodesorption can also be quantified by molecular dynamics simulations. In a series of papers, both the

photodesorption mechanisms and yields have been investigated for various water isotopologues and for different ice temperatures (Andersson et al., 2006; Andersson & van Dishoeck, 2008; Arasa et al., 2010, 2011; Koning et al., 2013). The incident UV photon is absorbed by a water molecule in the ice, which dissociates into H + OH, with both fragments having excess energy. The outcome depends on the ice layer in which the UV photon is absorbed; only the top layers actively participate in desorption. Intact water molecules are released to the gas with a yield of  $\sim 5 \times 10^{-4}$  per incident photon through two processes: (i) recombination of the H + OH fragments followed by escape of the energetic newly formed water molecule; and (ii) kick-out of a neighbouring water molecule by the energetic H atom produced by photodissociation. In contrast with molecules such as CO, the excited water molecule does not directly desorb since all UV absorptions immediately lead to dissociation of the molecule. Koning et al. (2013) have investigated the yields for the various processes using all combinations of water and its isotopologues, i.e., HDO and D<sub>2</sub>O in H<sub>2</sub>O ice but also H<sub>2</sub>O in D<sub>2</sub>O ice. Isotope-selective effects are found for the various combinations because of the different masses of the fragments. For example, if a D atom is created upon photodissociation, it is more effective in kicking out a neighbouring molecule than an H atom because of more efficient momentum transfer.

We here summarise the calculated photodesorption efficiencies as a function of ice layer and ice temperature for the astrophysically relevant cases of H<sub>2</sub>O, HDO, and D<sub>2</sub>O in H<sub>2</sub>O ice, and we provide convenient fitting formulae with depth into the ice for use in astrochemical models. Subsequently the isotope-selective effects are quantified. Specifically, the more effective desorption of atomic H compared with atomic D will result in enrichment of D in the ice. Furthermore, the extent to which the photodesorption process can affect the ortho/para ratio of H<sub>2</sub>O in the gas is briefly discussed. Sect. 2 briefly summarises the computational methods that were used in previous papers. Sect. 3 lists the results for the isotope selective processes for each ice layer at ice temperatures from 10–90 K. Sect. 4 summarises the conclusions and astrophysical implications.

## 2. Methods

Our methods have been explained in detail in our previous studies (Andersson et al., 2006; Andersson & van Dishoeck, 2008; Arasa et al., 2010, 2011; Koning et al., 2013), and are based on classical Molecular Dynamics (MD) methods (Allen & Tildesley, 1987), in which the atoms and molecules in water ice move according to Newton's equations based on analytical potentials of the interactions. In brief, an amorphous ice consisting of 480 water molecules is built for a certain ice temperature, and one water molecule (X<sub>2</sub>O (X=H or D) or HDO) is then randomly selected to be dissociated by UV radiation. Our previous studies show that only the photoexcited molecules that are initially located in the top four monolayers lead to photodesorption processes.

The dynamics of the photodissociation fragments are subsequently followed in the ice over a timescale of a few picoseconds until one of six possible outcomes is reached (see Fig. 1): (1) X atom desorption while OX stays trapped in the ice, (2) OX radical desorption while X stays trapped in the ice, (3) both X and OX desorb, (4) trapping of both X and OX photofragments, (5) recombination of the X atom and the OX radical to form X<sub>2</sub>O which in the end desorbs, or (6) stays trapped in the ice. When HDO is dissociated the outcome channels are the same, but the recombination of the photofragments leads to HDO that

either desorbs (5), or stays trapped in the ice (6). Moreover, there is a parallel outcome to any of the 6 processes by which one of the surrounding H<sub>2</sub>O molecules from the ice can also desorb, initiated by the same UV photon. This mechanism is called “kick-out” (Andersson et al., 2006; Andersson & van Dishoeck, 2008), and it takes place when the energetic X atom produced by photodissociation kicks a surrounding molecule from the ice by transfer of its momentum, and eventually the kicked out H<sub>2</sub>O molecule desorbs from the ice.

The calculations are repeated  $\geq 6000$  times per species and per monolayer so that the statistics (i.e., probabilities) of the different outcomes are determined as a function of the initial position of the molecule in the ice. The water molecules are excited with UV radiation in the 7.5–9.5 eV (1650–1300 Å) range, corresponding to the first electronic absorption band of water ice. Throughout the paper, \* indicates the molecule that is photodissociated in the ice. For the case of HDO\* there are two possible outcomes: HOD\* indicates photodissociation to H + OD whereas DOH\* denotes dissociation to D + OH; HDO and HDO\* denote the generic case.

In our previous studies (Arasa et al., 2010, 2011) we calculated the total photodesorption yields of OX and X<sub>2</sub>O per incident photon, and compared the results with the available experimental data (Öberg et al., 2009). For our model D<sub>2</sub>O ice, the photodesorption yield agrees well with the experimental photodesorption yield at low ice temperatures within the  $\sim 60\%$  experimental uncertainties (Arasa et al., 2011). At higher ice temperatures, the experimental photodesorption yield is larger than the calculated value due to long timescale thermal effects that cannot be considered in our simulations. The experimental OX/H<sub>2</sub>O ratio is however on average a factor of 2 to 3 lower than the calculated value for H<sub>2</sub>O ice (see Table II in Arasa et al. 2010), and D<sub>2</sub>O ice, respectively. In a different kind of experiment, Yabushita et al. (2009) measured the final translational and rotational energies of the kicked out H<sub>2</sub>O molecules for  $\nu = 0$  at  $T_{\text{ice}} = 90$  K; their values match well with our calculated ones at  $T_{\text{ice}} = 90$  K and provide proof for the importance of the kick-out mechanism (Arasa et al., 2010). These comparisons between models and data provide confidence in the accuracy of our simulations. For isotope selective processes, only relative trends are considered, which should have less uncertainty than the absolute values.

### 3. Results and discussion

In this section we summarize the probabilities of all the photodesorption events after H<sub>2</sub>O\*, HDO\* and D<sub>2</sub>O\* photodissociation in amorphous H<sub>2</sub>O ice at relevant temperatures,  $T_{\text{ice}} = 10$  and 90 K. The supplementary material contains the probabilities at  $T_{\text{ice}} = 20, 30$ , and 60 K for the top four monolayers (ML) of the ice surface. These results have been reported and extensively discussed elsewhere (Koning et al., 2013), but they are presented here in tabulated form for each of the top four monolayers so that the data can be used in astrochemical models. In addition, we provide fitting formulae to these tables in Table 1 (see Appendix B for more details). All tables can be downloaded from <http://www.strw.leidenuniv.nl/~ewine/photo>.

A total of 6000 trajectories per monolayer have been simulated for photoexcitation of H<sub>2</sub>O and D<sub>2</sub>O. For photoexcitation of HDO,  $\geq 6000$  trajectories were run per monolayer. The photodesorption probabilities are calculated for each event as  $p_i = N_i/N_{\text{total}}$ ,  $i$  = photodesorption event,  $N_i$  = number of trajectories that lead to the photodesorption event  $i$ ,  $N_{\text{total}}$  = total number of trajectories per monolayer  $j$ , the average is taken over

**Table 1.** Fitting functions and best-fit values of the parameters  $a$ ,  $b$ ,  $c$ , and ML, ML being the number of the monolayer.

Species	$a$	$b$	$c$
$X_{\text{des}} + OY_{\text{trapped}}$			
$P(\text{ML}) = (a\text{ML}) \exp(-b\text{ML})$			
H <sub>2</sub> O	2.02	0.858	...
D <sub>2</sub> O	2.06	0.906	...
HOD	2.04	0.867	...
DOH	2.17	0.942	...
$X_{\text{trapped}} + OY_{\text{des}}$			
$P(\text{ML}) = (a\text{ML}) \exp(-b(\text{ML} - c)^2)$			
H <sub>2</sub> O	4.28	0.0485	-10.9
D <sub>2</sub> O	0.0141	0.377	0.264
HOD	0.00359	0.635	0.990
DOH	0.0114	0.450	0.751
$X_{\text{des}} + OY_{\text{des}}$			
$P(\text{ML}) = (a\text{ML}) \exp(-b(\text{ML} - c)^2)$			
H <sub>2</sub> O	0.0710	0.210	-1.77
D <sub>2</sub> O	0.0638	0.554	0.269
HOD	0.0463	0.383	-0.383
DOH	0.0659	0.484	0.0538
$X_{\text{trapped}} + OY_{\text{trapped}}$			
$P(\text{ML}) = a(1 - \exp(-b\text{ML}^c))$			
H <sub>2</sub> O	0.516	0.149	1.48
D <sub>2</sub> O	0.458	0.163	1.57
HOD	0.595	0.125	1.60
DOH	0.574	0.111	1.89
$XYO_{\text{direct}}$			
$P(\text{ML}) = (a\text{ML}) \exp(-b(\text{ML} - c)^2)$			
H <sub>2</sub> O	0.390	0.0819	-6.04
D <sub>2</sub> O	0.00523	0.470	0.967
HOD	0.00428	0.604	1.12
DOH	0.00494	0.570	1.11
$H_2O_{\text{kicked}}$			
$P(\text{ML}) = (a\text{ML}) \exp(-b(\text{ML} - c)^2)$			
H <sub>2</sub> O	0.00320	0.210	0.882
D <sub>2</sub> O	0.00865	0.452	1.83
HOD	0.00465	0.193	0.419
DOH	0.00949	0.180	0.519

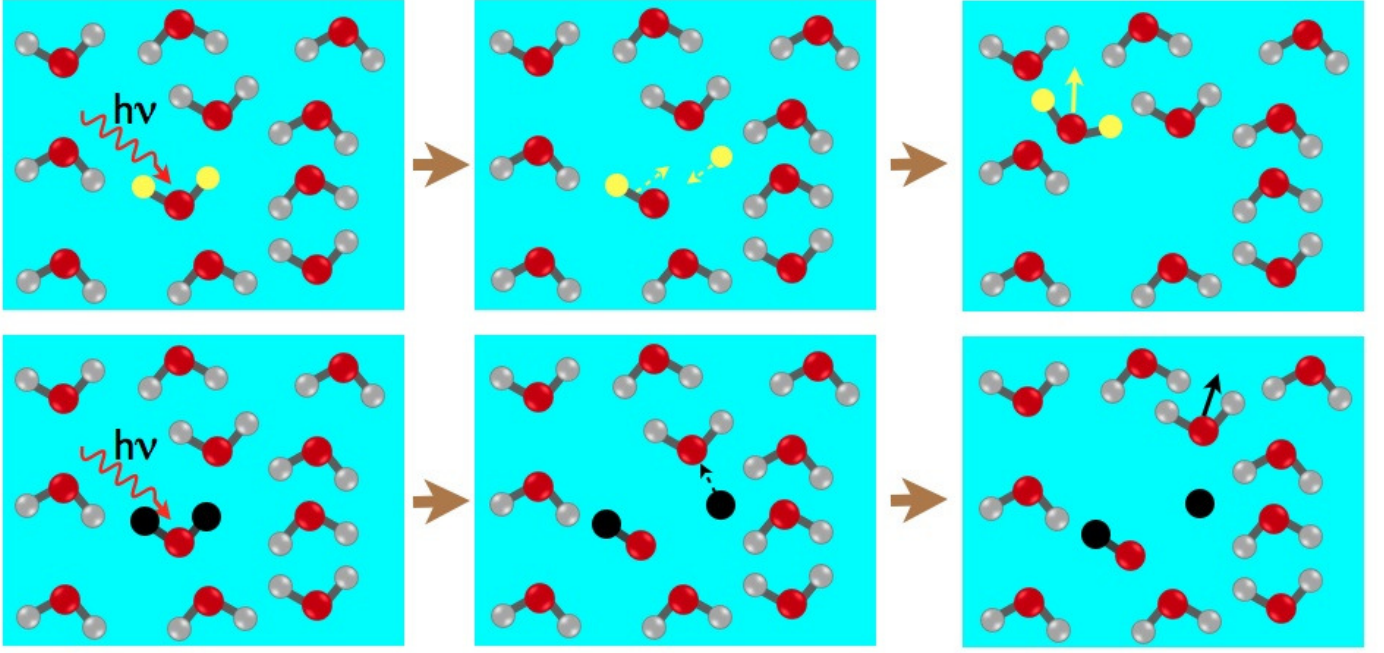
**Notes.** The species XOY denotes photoexcitation of XOY leading to dissociation into X + OY. The units of the parameters,  $a$ ,  $b$ , and  $c$ , depend on the fitting function which itself describes a probability and thus has no physical units. For outcome (1),  $X_{\text{des}} + OY_{\text{trapped}}$ ,  $a$  and  $b$  have units of ML<sup>-1</sup>, where ML is monolayer number. For outcome (4),  $X_{\text{trapped}} + OY_{\text{trapped}}$ ,  $a$  has no units and  $b$  has units of ML<sup>-c</sup>. For all other outcomes,  $a$  has units of ML<sup>-1</sup>,  $b$  has units of ML<sup>-2</sup>, and  $c$  has units of ML.

the top four monolayers (e.g.,  $\langle p_i \rangle = \sum_{j=1}^4 p_i^j / 4$ ). All the probabilities are given per *absorbed* (rather than incident) UV photon. The averaged photodesorption probabilities calculated per incident UV photon are on average a factor of  $\sim 0.03$  smaller than those given per absorbed UV photon for processes (1), (2), (3) and (5) (see numbering of outcomes in Sect. 2 of this paper and the description in Arasa et al., 2010, 2011).

Above 100 K, thermal desorption takes over as the main desorption mechanism in interstellar space, so our results are not relevant for the interpretation of data on warm water.

#### 3.1. X (X=H,D) atom photodesorption probabilities

The X atom photodesorption event is the dominant process when the X<sub>2</sub>O photodissociated molecule is located in the top four monolayers of the ice surface (see Table 2 and Andersson et al.,



**Fig. 1.** Top: A H<sub>2</sub>O molecule (with H atoms in yellow) surrounded by water molecules absorbs an UV photon and dissociates into H and OH. H and OH recombine to form H<sub>2</sub>O that eventually desorbs from the ice surface via the direct mechanism. Bottom: A D<sub>2</sub>O molecule (with D atoms in black) surrounded by water molecules absorbs an UV photon and dissociates into D and OD. The heavier D atom transfers its momentum to one of the surrounding H<sub>2</sub>O molecules that desorbs from the ice via the “kick-out” mechanism.

2006; Andersson & van Dishoeck, 2008). The X atom photodesorption probabilities are  $> 90\%$  in the top monolayer (ML1) and decrease with depth to  $\sim 30\%$  in ML4 because in the top MLs the X atoms can easily find their way to escape from the surface, whereas deeper in the ice there are other molecules that prevent their desorption.

In general, H atom photodesorption probabilities are somewhat larger following H<sub>2</sub>O dissociation than those for D atom photodesorption after D<sub>2</sub>O photodissociation in H<sub>2</sub>O ice, especially if initially the molecules are located in the third and fourth monolayers. The same trend is observed after HOD\*/DOH\* photodissociation in H<sub>2</sub>O ice: H atom photodesorption probabilities are larger than those for the D atom. This is because of isotope mass effects: the H atom is lighter than the D atom and therefore transfers less energy to nearby water molecules, so that it can travel more easily from the bottom to the top of the surface and eventually desorb to the gas phase (Arasa et al., 2011; Koning et al., 2013).

The average H/D photodesorption ratios following H<sub>2</sub>O\* and D<sub>2</sub>O\* photodissociation and the average H/D ratios following HDO\* photodissociation over the top four MLs are summarized in Table 3 versus ice temperature. The photodesorption ratio  $\frac{H_{des}}{D_{des}}$  after HDO photodissociation have been calculated (as was done for the photodesorption  $\frac{OD_{des}}{OH_{des}}$  in Table IV in Koning et al. 2013) taking into account the probabilities in each ML,  $i$ , where  $P_{D_{des}}(i, T_{ice})$  is the D photodesorption probability after DOH\* photodissociation into D and OH, and  $P_{H_{des}}(i, T_{ice})$  is the H photodesorption probability after HOD\* photodissociation into H and OD, and the branching ratios  $(\frac{H+OD}{D+OH})$  in the ice in each ML,  $i$ , and temperature  $(\beta(i, T_{ice}))$ . The effect is typically a factor of 1.1 when comparing H<sub>2</sub>O and D<sub>2</sub>O, and when comparing HOD in H<sub>2</sub>O and DOH in H<sub>2</sub>O.

**Table 3.** Average photodesorption ratio  $\frac{H_{des}}{D_{des}}$  following H<sub>2</sub>O\*, D<sub>2</sub>O\* and HDO\* photodissociation. For HDO\*, the last column shows the result without considering the effect of branching in the photodissociation.

$T_{ice}$ (K)	H <sub>2</sub> O* and D <sub>2</sub> O*	$\frac{H_{des}}{D_{des}}$	HDO*
10	$1.07 \pm 0.01$	$2.36 \pm 0.01$	$1.01 \pm 0.01$
20	$1.06 \pm 0.01$	$2.43 \pm 0.02$	$1.19 \pm 0.01$
30	$1.06 \pm 0.01$	$2.45 \pm 0.02$	$1.08 \pm 0.01$
60	$1.06 \pm 0.01$	$2.55 \pm 0.02$	$1.11 \pm 0.01$
90	$1.10 \pm 0.01$	$2.47 \pm 0.02$	$1.08 \pm 0.01$

For HDO\* photodissociation, there is another effect at work, namely that the probability of dissociation into H + OD is larger than that into D + OH, a process that is well known in the gas phase. For gas-phase photodissociation, the OD/OH branching ratio is about 3.1 (van Harreveld & van Hemert, 2012). In the ice, this effect is less, but still at the factor of two level. The HDO\* photodissociation branching ratios  $(\frac{H+OD}{D+OH})$  in the ice in each ML,  $i$ , and temperature  $(\beta(i, T_{ice}))$  were calculated and listed in Table IV in Koning et al. (2013). The average branching ratio taken over all ice temperatures is about 2.2 in favor of H + OD. Thus the branching ratio upon HDO photodissociation in the ice reinforces the H/D desorption ratio leading to a ratio more than a factor of two larger than the ratio of H and D present in the ice (calculated from Tables 2 and A.1). Therefore, all photons that arrive in the ice and are absorbed in the top layers of the ice surface lead to an enrichment in D atoms relative to H atoms in the ice mantles of dust particles.

**Table 2.** X atom photodesorption probabilities at  $T_{\text{ice}} = 10$  K (top) and 90 K (bottom) resulting from photoexcitation of a X<sub>2</sub>O (X=H,D) or XOY (HOD or DOH) molecule present in a specific monolayer of H<sub>2</sub>O ice.

ML	H <sub>des</sub> /H <sub>2</sub> O*	D <sub>des</sub> /D <sub>2</sub> O*	H <sub>des</sub> /HOD*	D <sub>des</sub> /DOH*
$T_{\text{ice}} = 10$ K				
1	$0.920 \pm 3.9 \times 10^{-3}$	$0.900 \pm 4.0 \times 10^{-3}$	$0.911 \pm 2.8 \times 10^{-3}$	$0.912 \pm 3.5 \times 10^{-3}$
2	$0.700 \pm 5.8 \times 10^{-3}$	$0.681 \pm 6.0 \times 10^{-3}$	$0.751 \pm 4.3 \times 10^{-3}$	$0.710 \pm 5.6 \times 10^{-3}$
3	$0.510 \pm 6.0 \times 10^{-3}$	$0.488 \pm 6.3 \times 10^{-3}$	$0.521 \pm 4.9 \times 10^{-3}$	$0.472 \pm 6.0 \times 10^{-3}$
4	$0.340 \pm 5.2 \times 10^{-3}$	$0.231 \pm 5.4 \times 10^{-3}$	$0.259 \pm 4.3 \times 10^{-3}$	$0.190 \pm 4.7 \times 10^{-3}$
$\langle \text{MLs} \rangle$	$0.617 \pm 3.2 \times 10^{-3}$	$0.575 \pm 3.2 \times 10^{-3}$	$0.610 \pm 4.1 \times 10^{-3}$	$0.572 \pm 5.0 \times 10^{-3}$
$T_{\text{ice}} = 90$ K				
1	$0.922 \pm 3.9 \times 10^{-3}$	$0.891 \pm 4.0 \times 10^{-3}$	$0.910 \pm 3.4 \times 10^{-3}$	$0.910 \pm 3.5 \times 10^{-3}$
2	$0.763 \pm 5.3 \times 10^{-3}$	$0.682 \pm 6.0 \times 10^{-3}$	$0.660 \pm 6.0 \times 10^{-3}$	$0.630 \pm 6.0 \times 10^{-3}$
3	$0.425 \pm 5.6 \times 10^{-3}$	$0.407 \pm 6.3 \times 10^{-3}$	$0.460 \pm 6.3 \times 10^{-3}$	$0.380 \pm 6.0 \times 10^{-3}$
4	$0.320 \pm 3.4 \times 10^{-3}$	$0.231 \pm 5.4 \times 10^{-3}$	$0.320 \pm 6.0 \times 10^{-3}$	$0.250 \pm 5.3 \times 10^{-3}$
$\langle \text{MLs} \rangle$	$0.608 \pm 3.2 \times 10^{-3}$	$0.553 \pm 3.2 \times 10^{-3}$	$0.589 \pm 5.5 \times 10^{-3}$	$0.543 \pm 5.0 \times 10^{-3}$

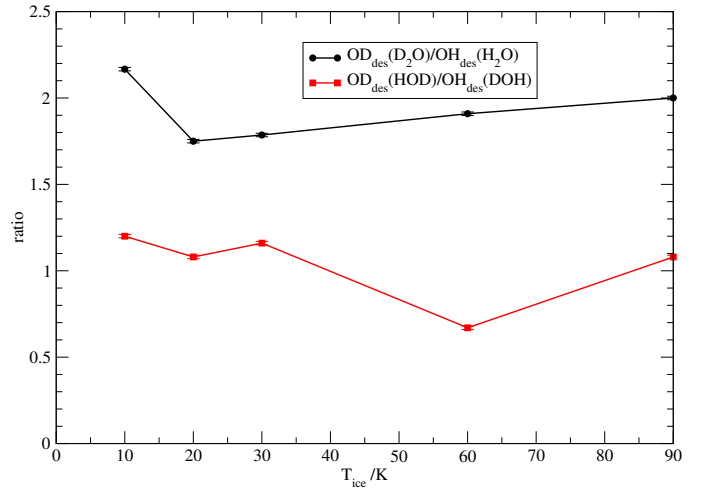
**Notes.** ML1 is the top monolayer and  $\langle \text{MLs} \rangle$  denotes the average over the top four monolayers.

### 3.2. OX (X=H,D) radical photodesorption probabilities

The second most important photodesorption event is OX photodesorption with probabilities of typically a few % in the top layers. In Table 4 the OX radical photodesorption probabilities for the top four monolayers together with the average values taken over the top four MLs are summarised at  $T_{\text{ice}} = 10$  K and 90 K. It is seen that the photodesorption probabilities drop sharply from ML2 to ML3. This feature has been explained in detail previously (Andersson et al., 2006; Andersson & van Dishoeck, 2008; Koning et al., 2013). Although OX is quite mobile in the top ice layer (ML1), it loses much of its energy in lower layers to other H<sub>2</sub>O molecules in the ice, preventing desorption of OX.

The OH photodesorption probabilities after DOH\* photodissociation and OD photodesorption probabilities after D<sub>2</sub>O\* photodissociation are larger than those for OD after HOD\* photodissociation and OH after H<sub>2</sub>O\* in H<sub>2</sub>O ice. This result is expected because the OD radicals have higher kinetic energies than the OH radicals upon the initial D<sub>2</sub>O and H<sub>2</sub>O photodissociation, respectively. In contrast, the OH radicals have higher kinetic energy than the OD radicals upon HDO\* photodissociation. These findings follow from applying energy and momentum conservation ( $p_{\text{OX}} + p_{\text{X}} = 0$ , and  $E_{\text{OX}} + E_{\text{X}} = \Delta E$ , with  $\Delta E$  being the initial available energy;  $\Delta E = E_{\text{exc}} - E_{\text{diss}}(\text{X}_2\text{O})$ , see Arasa et al. 2011). For X<sub>2</sub>O photodissociation at the same excitation energy  $E_{\text{exc}}$  it is seen that if  $E_{\text{OX}} = \Delta E / (1 + m_{\text{OX}}/m_{\text{X}})$ , the initial energy of OD should be a factor 1.8 larger than that for OH. In the case of X<sup>(1)</sup>OX<sup>(2)</sup> photodissociation into OX<sup>(2)</sup> + X<sup>(1)</sup>, and applying the same conservation rules we can expect that the energy of the OX<sup>(2)</sup> radical formed upon DOX<sup>(2)</sup> photodissociation is about a factor of two larger than the OX<sup>(2)</sup> radical formed upon HOX<sup>(2)</sup> photodissociation (see Table III in Koning et al. 2013).

The average OD/OH photodesorption ratios following HDO\* photodissociation over the top four MLs are summarised in Fig. 2. The photodesorption ratio OD<sub>des</sub>/OH<sub>des</sub> after HOD\* and DOH\* photodissociation have been calculated taking account the probabilities in each ML,  $i$ , where  $P_{\text{OD}_{\text{des}}}(i, T_{\text{ice}})$  is the OD photodesorption probability after HOD\* photodissociation into H and OD, and  $P_{\text{OH}_{\text{des}}}(i, T_{\text{ice}})$  is the OH photodesorption probability after DOH\* photodissociation into D and OH, and the photodissociation branching ratios H+OD/D+OH in the ice in each ML,  $i$ , and temperature ( $\beta(i, T_{\text{ice}})$ , see Table IV in Koning et al. 2013). These values are seen to be close to unity and to be mostly



**Fig. 2.**  $\text{OD}_{\text{des}}(\text{D}_2\text{O})/\text{OH}_{\text{des}}(\text{H}_2\text{O})$  ratios and  $\text{OD}_{\text{des}}(\text{HOD})/\text{OH}_{\text{des}}(\text{DOH})$  ratios averaged over the top four monolayers following HDO dissociation in H<sub>2</sub>O ice versus ice temperature  $T_{\text{ice}}$  (based on Koning et al., 2013).

constant with ice temperature. The reason that the OD<sub>des</sub>/OH<sub>des</sub> ratios are close to unity for HDO\* is that a cancellation of the two effects occurs. On the one hand, the OH desorption is more efficient than that of OD, as explained above. However, this is offset by the OD/OH branching ratio of about 2.2 in the HDO\* photodissociation, as discussed in Sect. 3.1. As a result, the average photodesorption ratio OD<sub>des</sub>/OH<sub>des</sub> upon HDO\* photodissociation is about 1.0 with a standard deviation of 0.2 in the ice (Koning et al., 2013). Note that this result is very different from pure gas-phase chemistry, where OD is produced a factor of three more rapidly than OH by photodissociation of gaseous HDO.

Fig. 2 also contains the average photodesorption ratios between OD desorption following D<sub>2</sub>O photodissociation in H<sub>2</sub>O ice and OH desorption following H<sub>2</sub>O photodissociation in H<sub>2</sub>O ice for all ice temperatures considered. On average, the  $\text{OD}_{\text{des}}(\text{D}_2\text{O})/\text{OH}_{\text{des}}(\text{H}_2\text{O})$  ratio is about 1.9.

**Table 4.** OX radical photodesorption probabilities at  $T_{\text{ice}} = 10$  K (top) and 90 K (bottom) resulting from photoexcitation of a X<sub>2</sub>O (X=H,D) or XOY (HOD or DOH) molecule present in a specific monolayer of H<sub>2</sub>O ice.

ML	OH <sub>des</sub> /H <sub>2</sub> O*	OD <sub>des</sub> /D <sub>2</sub> O*	OD <sub>des</sub> /HOD*	OH <sub>des</sub> /DOH*
$T_{\text{ice}} = 10$ K				
1	$0.022 \pm 1.9 \times 10^{-3}$	$0.065 \pm 3.2 \times 10^{-3}$	$0.044 \pm 2.0 \times 10^{-3}$	$0.071 \pm 3.2 \times 10^{-3}$
2	$0.021 \pm 1.8 \times 10^{-3}$	$0.035 \pm 2.8 \times 10^{-3}$	$0.014 \pm 1.2 \times 10^{-3}$	$0.033 \pm 2.2 \times 10^{-3}$
3	$(3.0 \pm 0.7) \times 10^{-3}$	$(4.0 \pm 0.9) \times 10^{-3}$	$(9.6 \pm 3.0) \times 10^{-4}$	$(5.1 \pm 0.9) \times 10^{-3}$
4	0.00	0.00	$(1.9 \pm 1.3) \times 10^{-4}$	0.00
⟨MLs⟩	$0.012 \pm 6.9 \times 10^{-4}$	$0.026 \pm 1.0 \times 10^{-3}$	$0.015 \pm 1.2 \times 10^{-3}$	$0.027 \pm 2.0 \times 10^{-3}$
$T_{\text{ice}} = 90$ K				
1	$0.036 \pm 1.8 \times 10^{-3}$	$0.065 \pm 3.2 \times 10^{-3}$	$0.016 \pm 1.6 \times 10^{-3}$	$0.039 \pm 2.4 \times 10^{-3}$
2	$0.025 \pm 1.9 \times 10^{-3}$	$0.050 \pm 2.8 \times 10^{-3}$	$0.023 \pm 1.9 \times 10^{-3}$	$0.045 \pm 2.6 \times 10^{-3}$
3	$(2.0 \pm 0.7) \times 10^{-3}$	$(5.0 \pm 0.9) \times 10^{-3}$	$(3.8 \pm 0.8) \times 10^{-3}$	$(7.4 \pm 0.1) \times 10^{-3}$
4	0.00	$(3.3 \pm 2.4) \times 10^{-4}$	$(1.6 \pm 1.6) \times 10^{-4}$	$(3.0 \pm 2.2) \times 10^{-4}$
⟨MLs⟩	$0.015 \pm 7.9 \times 10^{-4}$	$0.030 \pm 4.0 \times 10^{-4}$	$0.011 \pm 1.3 \times 10^{-3}$	$0.023 \pm 2.0 \times 10^{-3}$

**Notes.** ML1 is the top monolayer and ⟨MLs⟩ denotes the average over the top four monolayers.

### 3.3. X<sub>2</sub>O (X=H,D) and HDO molecule photodesorption probabilities

X<sub>2</sub>O photodesorption is the third most important photodesorption event after H<sub>2</sub>O\* and D<sub>2</sub>O\* photodissociation with probabilities of  $\sim 0.5\%$  in the top layers. As discussed above, two processes can occur: direct desorption of the recombined H<sub>2</sub>O or D<sub>2</sub>O molecule or through the kick of the energetic X atom to a surrounding H<sub>2</sub>O molecule from the ice. When HOD\* dissociates it can also lead to the recombination of the H and OD fragments that eventually desorb as HDO, or the H atom can kick out the surrounding H<sub>2</sub>O molecules from the ice. Similarly, when DOH\* dissociates the photofragments can also recombine and desorb, or the D atom can kick out a H<sub>2</sub>O molecule. In Table 5 the probabilities of both mechanisms for the top four MLs and its average for the four different scenarios are reported at  $T_{\text{ice}} = 10$  K and 90 K. Again, isotope effects are rather noticeable if X<sub>2</sub>O desorbs from the ice surface through the kick-out mechanism (Koning et al., 2013) after the X atom transfers its momentum to the surrounding H<sub>2</sub>O molecules. Since the D atom is heavier than the H atom it can transfer its momentum more easily to a surrounding H<sub>2</sub>O molecule, leading to higher kick-out photodesorption probabilities, and in consequence to higher total X<sub>2</sub>O and HOD photodesorption probabilities (Arasa et al., 2011; Koning et al., 2013). As demonstrated in Appendix C, this process does however not lead to fractionation of HDO/H<sub>2</sub>O in the gas in H<sub>2</sub>O-dominated ices, i.e., the gas-phase HDO/H<sub>2</sub>O ratio produced by photodesorption reflects that originally in the ice.

### 3.4. Ortho/para ratio of photodesorbed H<sub>2</sub>O

Observations of water in space also provide constraints on the ortho/para ratio of water in the gas. In diffuse interstellar clouds and high temperature shocks, these measurements are consistent with a ratio of 3, as expected from gas phase formation of water (e.g., Flagey et al., 2013; Herczeg et al., 2012; Emprechtinger et al., 2013). In contrast, the ortho/para ratio of water in protoplanetary disks and in photon-dominated regions, such as the Orion Bar, is found to be much lower, with values  $< 1$  (Hogerheijde et al., 2011; Choi et al., 2014). In these regions, the water results from photodesorption of water ice, not from thermal sublimation as is the case for comets (Mumma & Charnley, 2011). The

question therefore arises to what extent the process of photodesorption preserves the ortho/para ratio that was present in the ice.

In the direct photodesorption mechanism, the H-OH bond is broken and then reformed, so the ortho/para ratio is basically reset to the statistical value of 3. For the kick-out mechanism, however, the original ortho/para ratio in the ice should be preserved. The relative importance of the direct and kick-out mechanisms depends on the ice monolayer and to a lesser extent on the ice temperature. Table 5 summarises the average values over the first four monolayers of the two options. It is seen that the direct mechanism is typically a factor 1.3 more efficient than the kick-out mechanism for H<sub>2</sub>O. Thus, if the ortho/para ratio in the ice were low, say  $\leq 0.5$ , the gas-phase ortho/para ratio of the photodesorbed H<sub>2</sub>O would be significantly higher, at least 2, due to the photodesorption process.

Low ortho/para ratios in the ice could occur if the ratio had equilibrated to the grain temperature of typically 10–30 K, a process that in itself is still poorly understood (e.g., Dulieu, 2011). The main message here is that due to the partial reset during the photodesorption process, the observed gaseous ortho/para ratio only partially reflects the original ice ratio.

## 4. Conclusions

This paper has investigated the deuterium fractionation that can occur when OD, HDO or D<sub>2</sub>O molecules are liberated from a H<sub>2</sub>O-rich ice following a photodesorption event at a microscopic level. The results presented in Sect. 3 lead to the following conclusions:

1. No isotope fractionation occurs in the photodesorption of HDO and H<sub>2</sub>O from a mixed ice, even though kick-out by D is more efficient. The ratio of photodesorbed HDO and H<sub>2</sub>O is equal to the ratio of HDO and H<sub>2</sub>O in the ice.
2. No isotope fractionation occurs in the photodesorption of OD and OH upon HDO photodissociation in H<sub>2</sub>O ice. The ratio of photodesorbed OD and OH is equal to the HDO/H<sub>2</sub>O ratio in the ice within 20%.
3. The ratio of photodesorbed H and D will be more than twice the ratio of H and D present in the ice in HDO and H<sub>2</sub>O, because photo-excitation of HDO leads preferentially to desorption of H. Therefore, given enough time and enough photons photoprocessing of the outer layers of icy mantles could lead to enrichment in D relative to H.



**Table 5.** Total X<sub>2</sub>O (X=H,D) or XOY (HOD or DOH) photodesorption probabilities at  $T_{\text{ice}} = 10$  K (top) and 90 K (bottom) per monolayer due to the direct and the kick-out mechanism for X<sub>2</sub>O and XOY photodissociation in H<sub>2</sub>O ice.

ML	H <sub>2</sub> O*		D <sub>2</sub> O*		HOD*		DOH*	
	H <sub>2</sub> O <sub>direct</sub>	H <sub>2</sub> O <sub>kicked</sub>	D <sub>2</sub> O <sub>direct</sub>	H <sub>2</sub> O <sub>kicked</sub>	HOD <sub>direct</sub>	H <sub>2</sub> O <sub>kicked</sub>	DOH <sub>direct</sub>	H <sub>2</sub> O <sub>kicked</sub>
$T_{\text{ice}} = 10$ K								
1	5.7 ± 1.1	0.5 ± 0.2	6.8 ± 1.1	12.8 ± 1.5	5.6 ± 0.7	2.2 ± 0.5	5.3 ± 0.9	2.9 ± 0.7
2	7.4 ± 1.2	3.9 ± 1.3	9.2 ± 1.2	17.8 ± 1.7	5.1 ± 0.7	2.7 ± 0.5	7.0 ± 1.0	10.7 ± 1.3
3	1.7 ± 0.8	4.8 ± 1.1	1.0 ± 0.4	16.5 ± 1.6	2.9 ± 0.5	0.8 ± 0.3	1.6 ± 0.5	1.2 ± 0.4
4	0.00	0.6 ± 0.3	0.3 ± 0.2	4.3 ± 0.9	0.1 ± 0.1	1.2 ± 0.3	0.00	0.9 ± 0.4
⟨MLs⟩	3.7 ± 0.4	2.5 ± 0.2	4.3 ± 0.4	6.6 ± 0.5	3.4 ± 0.6	1.7 ± 0.4	3.5 ± 0.7	3.9 ± 0.8
$T_{\text{ice}} = 90$ K								
1	5.3 ± 1.1	1.4 ± 0.8	6.8 ± 0.1	12.8 ± 1.5	2.6 ± 0.6	5.4 ± 0.9	6.8 ± 0.1	15.6 ± 1.5
2	9.6 ± 1.3	3.9 ± 0.9	9.2 ± 1.2	17.8 ± 1.7	6.7 ± 1.1	12.2 ± 1.4	6.6 ± 1.0	21.8 ± 1.8
3	2.6 ± 0.8	5.2 ± 0.9	1.0 ± 0.4	16.5 ± 1.6	1.0 ± 0.4	7.2 ± 1.1	0.8 ± 0.3	21.0 ± 1.8
4	0.00	2.3 ± 0.7	0.3 ± 0.2	4.3 ± 0.9	0.3 ± 0.2	2.9 ± 0.7	1.2 ± 0.4	0.8 ± 0.3
⟨MLs⟩	4.4 ± 0.4	3.2 ± 0.4	4.3 ± 0.1	12.0 ± 1.0	2.7 ± 0.7	6.9 ± 1.0	3.9 ± 0.8	15.0 ± 1.5

**Notes.** All values should be multiplied by  $10^{-3}$ . The notation H<sub>2</sub>O\* and HOD\* denote photoexcitation of H<sub>2</sub>O and HDO in H<sub>2</sub>O ice in which the H atom resulting from photodissociation kicks out a neighbouring H<sub>2</sub>O molecule. The notation D<sub>2</sub>O\* and DOH\* indicate the analogous process in which the resulting D atom kicks out a neighbouring H<sub>2</sub>O molecule.

A possible mechanism for an increased ratio of HDO/H<sub>2</sub>O in the gas phase compared with that initially in the ice may therefore be prolonged photoprocessing of the outer layers of icy mantles of dust particles. Photo-excitation of HDO leads preferentially to desorption of H. As a consequence, the ratio of photodesorbed H and D will be greater than twice the ratio of H and D present in the ice. If photoprocessing occurs over a long enough time, the outer layers will develop a larger fraction of D, and, through grain surface reactions of D with OH, therefore also of HDO. So ultimately, the ratio of photodesorbed HDO and H<sub>2</sub>O can become enhanced due to this indirect process. Full gas-grain models that take the multi-layer structure of the ice into account are needed to investigate whether this is a plausible mechanism for isotope fractionation of HDO, since pure gas-phase processes can also lead to fractionation (e.g., Aikawa et al., 2012). If it is, then the relative abundance of HDO may also reflect the total fluence of UV radiation (product of UV flux and time) that the outer layers of icy mantles have been exposed to in the environment where they are observed.

Finally, we note that the ortho/para ratio of water in the gas does not directly reflect that in the ice but is partially reset by the photodesorption process.

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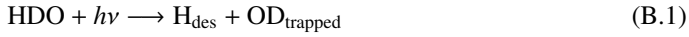
## Appendix A: Auxiliary tables

Tables A.1, A.2, and A.3 contain the probabilities for X atom desorption, OX desorption, and X<sub>2</sub>O and HDO desorption (X=H,D), respectively, at  $T_{\text{ice}} = 20, 30, \text{ and } 60 \text{ K}$ .

## Appendix B: Fitting formulae for photodesorption probabilities

Tables 2, 4, 5, A.1, A.2, and A.3 list the *total* probabilities for X desorption, OX desorption, and X<sub>2</sub>O and HDO desorption (X=H,D) following a dissociation event, as a function of both monolayer and ice temperature. These tables also give the average probabilities, over the top four monolayers, for each species. For use in astrochemical models, it is useful to know the probability (per monolayer) of *every* potential outcome, rather than the total probability for the desorption of each species. This is because, in full gas-grain models, one is also interested in the composition of the ice mantle, as well as the gas.

As discussed in the main body of this paper, there are six potential outcomes following a dissociation event which can lead to a change in composition of both the ice and gas. For example, for HDO which is dissociated into H + OD,



Eq. B.6 is the process known as “kick out” whereby a neighbouring H<sub>2</sub>O is ejected from the ice via momentum transfer from an excited photofragment. The probabilities of each of these events as a function of monolayer and ice temperature have been compiled from the raw data of the molecular dynamics simulations and are available in electronic form only. There is a seventh possibility in which the photofragments recombine to reform HDO which remains trapped in the ice. This process does not change the gas or ice composition and thus we have not listed the probabilities for this outcome here; however, these data are necessary if one is interested in extrapolating the probabilities to deeper monolayers,  $\text{ML} > 4$ .

To determine the desorption probabilities at temperatures and in monolayers outside of those tabulated, one can simply interpolate/extrapolate using, for example, cubic spline interpolation. However, when extrapolating to determine probabilities for deeper monolayers,  $\text{ML} > 4$ , one should take care to ensure that, deep into the ice mantle, the probabilities for outcomes B.1, B.2, B.3, B.5, and B.6 tend to 0, and the probability for outcome B.4 tends to  $1 - P_{\text{recom}}$ , where  $P_{\text{recom}}$  is the probability that the photofragments recombine to reform the molecule (which remains trapped in the ice). Deeper into the ice, desorption events become increasingly less probable and the most probable outcome becomes trapping of the photofragments (or the reformed molecule, following recombination). In addition, at very low coverage,  $\text{ML} < 1$ , the rates for all outcomes should tend to 0 as  $\text{ML} \rightarrow 0$ .

In Table 1 we present our fitting functions and corresponding best-fit parameters for the *temperature-averaged* probabilities per monolayer for each outcome. The probabilities are well fitted using a Gaussian-like function with the exception of the outcomes leading to trapping of the OY radical for which an

exponential-like function was found to be more appropriate for describing the asymptotic behaviour of the probabilities towards deeper monolayers ( $\gg 4$ ). In Fig. B.1 we present the probability per monolayer at each temperature and the temperature-averaged probabilities per monolayer along with the fitted functions for the example of DOH\*. The probabilities were fitted using the nonlinear least-squares (NLLS) Marquardt-Levenberg algorithm (Marquardt, 1963). The probabilities are a much stronger function of monolayer than temperature; hence, our decision to fit functions with respect to monolayer only.

For implementation in chemical models which adopt the rate equation method for describing the ice chemistry and gas-grain balance, the probabilities per monolayer should be multiplied by the rate of arrival of UV photons in the wavelength range 1650–1300 Å onto the grain surface times the absorption cross section of a UV photon by a grain-surface site (or molecule, in this case, HDO). The total desorption rate is then determined by integrating the desorption rate per monolayer over the total number of monolayers on the grain. The probabilities can be directly employed in stochastic chemical models in which the discrete nature of chemical reactions are taken into account (see, e.g., Cuppen & Herbst, 2007).

## Appendix C: Photodesorption and fractionation

This section investigates whether photodesorption ultimately also leads to fractionation of HDO/H<sub>2</sub>O in the gas. We can estimate the total photodesorption probability ratio between HDO and H<sub>2</sub>O by taking into account the direct and kicked out mechanism in both cases. The probability of HDO photodesorption through the direct mechanism is given by

$$P_{\text{HDO}}^{\text{direct}} = r_{\text{HDO}} P^{\text{direct}}(\text{HDO}^*) \quad (\text{C.1})$$

In Eq. C.1,  $P^{\text{direct}}(\text{HDO}^*)$  is the probability that upon photoexcitation of HDO (the generic case) the HDO recombines and desorbs directly. It can be approximately calculated using

$$P_{\text{direct}}(\text{HDO}^*) = \frac{2}{3} P_{\text{direct}}(\text{HOD}^*) + \frac{1}{3} P_{\text{direct}}(\text{DOH}^*) \quad (\text{C.2})$$

and  $r_{\text{HDO}}$  is the original HDO/H<sub>2</sub>O ratio in the ice (of the order of 0.01 or less as indicated by observations). In Eq. C.2, the probabilities on the right hand side are the probabilities for the direct mechanism for photodesorbing HDO averaged over the top four monolayers and presented in Tables 5 and A.3.

The probability of H<sub>2</sub>O photodesorption through the direct mechanism is given by

$$P_{\text{H}_2\text{O}}^{\text{direct}} = (1 - r_{\text{HDO}}) P_{\text{direct}}(\text{H}_2\text{O}^*) \quad (\text{C.3})$$

In Eq. C.3,  $P_{\text{direct}}(\text{H}_2\text{O}^*)$  is the probability that upon photoexcitation H<sub>2</sub>O recombines and desorbs directly. It can be obtained directly from Tables 5 and A.3.

As can be seen from Table 2 and after using Eq. C.2,  $P_{\text{direct}}(\text{HDO}^*)$  and  $P_{\text{direct}}(\text{H}_2\text{O}^*)$  are roughly the same. As a result

$$P_{\text{HDO}}^{\text{direct}} / P_{\text{H}_2\text{O}}^{\text{direct}} = r_{\text{HDO}} / (1 - r_{\text{HDO}}) = r_{\text{HDO}} / r_{\text{H}_2\text{O}} \quad (\text{C.4})$$

meaning that there is no isotope fractionation due to the direct mechanism.

Now consider the kick-out mechanism. The indirect probabilities can be written as follows:

$$P_{\text{HDO}}^{\text{KOdes}} = r_{\text{HDO}} \times r_{\text{HDO}} \times P_{\text{KO}}(\text{HDO}; \text{HDO}^*) + (1 - r_{\text{HDO}}) \times r_{\text{HDO}} \times P_{\text{KO}}(\text{HDO}; \text{H}_2\text{O}^*) \quad (\text{C.5})$$

**Table A.1.** X atom photodesorption probabilities at  $T_{\text{ice}} = 20$  K, 30 K, and 60 K resulting from photoexcitation of a X<sub>2</sub>O (X=H,D) or XOY (HOD or DOH) molecule present in a specific monolayer of H<sub>2</sub>O ice.

ML	H <sub>des</sub> /H <sub>2</sub> O*	D <sub>des</sub> /D <sub>2</sub> O*	H <sub>des</sub> /HOD*	D <sub>des</sub> /DOH*
$T_{\text{ice}} = 20$ K				
1	$0.865 \pm 4.0 \times 10^{-3}$	$0.898 \pm 3.9 \times 10^{-3}$	$0.913 \pm 3.6 \times 10^{-3}$	$0.907 \pm 3.7 \times 10^{-3}$
2	$0.741 \pm 5.9 \times 10^{-3}$	$0.665 \pm 6.1 \times 10^{-3}$	$0.661 \pm 6.0 \times 10^{-3}$	$0.626 \pm 6.0 \times 10^{-3}$
3	$0.422 \pm 6.1 \times 10^{-3}$	$0.401 \pm 6.3 \times 10^{-3}$	$0.472 \pm 6.3 \times 10^{-3}$	$0.419 \pm 6.0 \times 10^{-3}$
4	$0.261 \pm 4.9 \times 10^{-3}$	$0.190 \pm 5.1 \times 10^{-3}$	$0.194 \pm 5.0 \times 10^{-3}$	$0.137 \pm 4.2 \times 10^{-3}$
⟨MLs⟩	$0.572 \pm 3.2 \times 10^{-3}$	$0.538 \pm 3.2 \times 10^{-3}$	$0.559 \pm 5.3 \times 10^{-3}$	$0.527 \pm 5.0 \times 10^{-3}$
$T_{\text{ice}} = 30$ K				
1	$0.878 \pm 4.3 \times 10^{-3}$	$0.871 \pm 4.3 \times 10^{-3}$	$0.863 \pm 4.3 \times 10^{-3}$	$0.864 \pm 4.3 \times 10^{-3}$
2	$0.740 \pm 5.9 \times 10^{-3}$	$0.670 \pm 6.1 \times 10^{-3}$	$0.711 \pm 5.7 \times 10^{-3}$	$0.655 \pm 5.8 \times 10^{-3}$
3	$0.435 \pm 6.3 \times 10^{-3}$	$0.452 \pm 6.4 \times 10^{-3}$	$0.507 \pm 6.3 \times 10^{-3}$	$0.453 \pm 6.0 \times 10^{-3}$
4	$0.319 \pm 5.2 \times 10^{-3}$	$0.240 \pm 5.6 \times 10^{-3}$	$0.310 \pm 5.8 \times 10^{-3}$	$0.236 \pm 5.1 \times 10^{-3}$
⟨MLs⟩	$0.593 \pm 3.2 \times 10^{-3}$	$0.558 \pm 3.2 \times 10^{-3}$	$0.598 \pm 5.6 \times 10^{-3}$	$0.552 \pm 5.0 \times 10^{-3}$
$T_{\text{ice}} = 60$ K				
1	$0.872 \pm 4.3 \times 10^{-3}$	$0.872 \pm 4.3 \times 10^{-3}$	$0.904 \pm 3.7 \times 10^{-3}$	$0.883 \pm 4.0 \times 10^{-3}$
2	$0.711 \pm 6.0 \times 10^{-3}$	$0.706 \pm 5.9 \times 10^{-3}$	$0.696 \pm 5.8 \times 10^{-3}$	$0.667 \pm 5.8 \times 10^{-3}$
3	$0.404 \pm 6.3 \times 10^{-3}$	$0.350 \pm 6.2 \times 10^{-3}$	$0.376 \pm 6.1 \times 10^{-3}$	$0.298 \pm 5.6 \times 10^{-3}$
4	$0.302 \pm 5.9 \times 10^{-3}$	$0.230 \pm 5.4 \times 10^{-3}$	$0.306 \pm 5.7 \times 10^{-3}$	$0.204 \pm 4.8 \times 10^{-3}$
⟨MLs⟩	$0.572 \pm 3.2 \times 10^{-3}$	$0.539 \pm 3.2 \times 10^{-3}$	$0.570 \pm 5.4 \times 10^{-3}$	$0.513 \pm 5.0 \times 10^{-3}$

**Notes.** ML1 is the top monolayer and ⟨MLs⟩ denotes the average over the top four monolayers.

**Table A.2.** OX radical photodesorption probabilities at  $T_{\text{ice}} = 20$  K, 30 K, and 60 K resulting from photoexcitation of a X<sub>2</sub>O (X=H,D) or XOY (HOD or DOH) molecule present in a specific monolayer of H<sub>2</sub>O ice.

ML	OH <sub>des</sub> /H <sub>2</sub> O*	OD <sub>des</sub> /D <sub>2</sub> O*	OD <sub>des</sub> /HOD*	OH <sub>des</sub> /DOH*
$T_{\text{ice}} = 20$ K				
1	$0.026 \pm 2.7 \times 10^{-3}$	$0.043 \pm 2.6 \times 10^{-3}$	$0.028 \pm 2.1 \times 10^{-3}$	$0.055 \pm 2.9 \times 10^{-3}$
2	$0.021 \pm 2.7 \times 10^{-3}$	$0.038 \pm 2.5 \times 10^{-3}$	$0.022 \pm 1.8 \times 10^{-3}$	$0.049 \pm 2.7 \times 10^{-3}$
3	$(1.0 \pm 0.3) \times 10^{-3}$	$(1.7 \pm 0.5) \times 10^{-3}$	$(4.7 \pm 2.7) \times 10^{-4}$	$(1.7 \pm 0.5) \times 10^{-3}$
4	0.00	0.00	0.00	0.00
⟨MLs⟩	$0.012 \pm 7.0 \times 10^{-4}$	$0.021 \pm 9.1 \times 10^{-4}$	$0.012 \pm 1.4 \times 10^{-3}$	$0.026 \pm 2.0 \times 10^{-3}$
$T_{\text{ice}} = 30$ K				
1	$0.047 \pm 2.9 \times 10^{-3}$	$0.069 \pm 3.3 \times 10^{-3}$	$0.029 \pm 2.1 \times 10^{-3}$	$0.054 \pm 2.8 \times 10^{-3}$
2	$0.011 \pm 1.8 \times 10^{-3}$	$0.019 \pm 1.8 \times 10^{-3}$	$(5.5 \pm 0.9) \times 10^{-3}$	$0.015 \pm 1.5 \times 10^{-3}$
3	$(1.7 \pm 0.5) \times 10^{-4}$	$0.012 \pm 1.4 \times 10^{-3}$	$(3.8 \pm 0.8) \times 10^{-3}$	$(6.8 \pm 0.9) \times 10^{-3}$
4	0.00	$(1.7 \pm 1.7) \times 10^{-4}$	$(2.0 \pm 2.0) \times 10^{-4}$	0.00
⟨MLs⟩	$0.014 \pm 7.5 \times 10^{-4}$	$0.025 \pm 1.0 \times 10^{-3}$	$(9.5 \pm 1.2) \times 10^{-3}$	$0.019 \pm 2.0 \times 10^{-3}$
$T_{\text{ice}} = 60$ K				
1	$0.029 \pm 2.2 \times 10^{-3}$	$0.053 \pm 2.9 \times 10^{-3}$	$0.013 \pm 1.4 \times 10^{-3}$	$0.050 \pm 2.7 \times 10^{-3}$
2	$0.012 \pm 1.4 \times 10^{-3}$	$0.026 \pm 2.0 \times 10^{-3}$	$(7.0 \pm 1.0) \times 10^{-3}$	$0.020 \pm 1.7 \times 10^{-3}$
3	$(2.3 \pm 0.6) \times 10^{-3}$	$(6.0 \pm 0.1) \times 10^{-3}$	$(3.5 \pm 0.8) \times 10^{-3}$	$0.012 \pm 1.3 \times 10^{-3}$
4	0.00	0.00	0.00	0.00
⟨MLs⟩	$0.011 \pm 6.7 \times 10^{-4}$	$0.021 \pm 9.3 \times 10^{-4}$	$(5.9 \pm 1.0) \times 10^{-3}$	$0.020 \pm 2.0 \times 10^{-3}$

**Notes.** ML1 is the top monolayer and ⟨MLs⟩ denotes the average over the top four monolayers.

and

$$P_{\text{H}_2\text{O}}^{\text{KOdes}} = (1 - r_{\text{HDO}}) \times (1 - r_{\text{HDO}}) \times P_{\text{KO}}(\text{H}_2\text{O}; \text{H}_2\text{O}^*) + r_{\text{HDO}} \times (1 - r_{\text{HDO}}) \times P_{\text{KO}}(\text{H}_2\text{O}; \text{HDO}^*) \quad (\text{C.6})$$

In Eqs. C.5 and C.6,  $P_{\text{HXO}}^{\text{KOdes}}$  is the probability of desorption of HXO through the kick-out mechanism, where X is either H or D. Furthermore,  $P_{\text{KO}}(\text{HX}^1\text{O}; \text{HX}^2\text{O}^*)$  is the probability that HX<sup>1</sup>O is kicked out after photo-excitation of HX<sup>2</sup>O, where X<sup>1</sup> can either be H or D, and X<sup>2</sup> can also be H or D. As for the direct mechanism, we can approximately calculate

$P_{\text{KO}}(\text{HX}^1\text{O}; \text{HX}^2\text{O}^*)$  from

$$P_{\text{KO}}(\text{HX}^1\text{O}; \text{HDO}^*) = \frac{2}{3} P_{\text{KO}}(\text{HX}^1\text{O}; \text{HOD}^*) + \frac{1}{3} P_{\text{KO}}(\text{HX}^1\text{O}; \text{DOH}^*) \quad (\text{C.7})$$

The two quantities on the right hand side of Eq. C.7 have been tabulated for X<sup>1</sup> equal to H in Tables 5 and A.3.

Because we have only calculated probabilities that H<sub>2</sub>O is kicked out, we make the following approximations,

$$P_{\text{KO}}(\text{HDO}; \text{HOD}^*) = P_{\text{KO}}(\text{H}_2\text{O}; \text{HOD}^*) \quad (\text{C.8})$$

$$P_{\text{KO}}(\text{HDO}; \text{DOH}^*) = P_{\text{KO}}(\text{H}_2\text{O}; \text{DOH}^*) \quad (\text{C.9})$$

**Table A.3.** Total X<sub>2</sub>O (X=H,D) or XOY (HOD or DOH) photodesorption probabilities at  $T_{\text{ice}} = 20$  K, 30 K, and 60 K per monolayer due to the direct and the kick-out mechanism for X<sub>2</sub>O and XOY photodissociation in H<sub>2</sub>O ice.

ML	H <sub>2</sub> O*		D <sub>2</sub> O*		HOD*		DOH*	
	H <sub>2</sub> O <sub>direct</sub>	H <sub>2</sub> O <sub>kicked</sub>	D <sub>2</sub> O <sub>direct</sub>	H <sub>2</sub> O <sub>kicked</sub>	HOD <sub>direct</sub>	H <sub>2</sub> O <sub>kicked</sub>	DOH <sub>direct</sub>	H <sub>2</sub> O <sub>kicked</sub>
$T_{\text{ice}} = 20$ K								
1	7.8 ± 1.3	0.7 ± 0.4	3.3 ± 0.7	3.3 ± 0.7	4.5 ± 0.9	1.6 ± 0.5	3.7 ± 0.8	4.2 ± 0.8
2	5.5 ± 1.2	2.8 ± 0.9	7.1 ± 1.1	12.5 ± 1.4	7.1 ± 1.1	3.7 ± 0.8	5.9 ± 0.9	4.2 ± 0.8
3	0.3 ± 0.2	5.9 ± 1.1	1.7 ± 0.5	13.0 ± 1.5	1.0 ± 0.4	0.8 ± 0.4	0.3 ± 0.2	7.0 ± 1.0
4	0.2 ± 0.2	0.2 ± 0.2	0.00	14.0 ± 1.5	0.2 ± 0.2	3.9 ± 0.8	0.00	16.0 ± 1.5
⟨MLs⟩	3.4 ± 0.4	2.4 ± 0.3	3.0 ± 0.4	11.0 ± 0.7	3.2 ± 0.7	2.5 ± 0.6	2.5 ± 0.6	7.9 ± 1.1
$T_{\text{ice}} = 30$ K								
1	2.6 ± 0.9	4.2 ± 0.9	4.5 ± 0.9	0.5 ± 0.3	4.6 ± 0.9	1.0 ± 0.4	5.6 ± 0.9	0.2 ± 0.2
2	2.9 ± 1.1	5.7 ± 0.9	4.8 ± 0.9	17.0 ± 0.2	3.6 ± 0.8	2.9 ± 0.7	7.5 ± 0.1	14.4 ± 1.5
3	1.7 ± 1.2	10.0 ± 0.2	2.2 ± 0.6	10.6 ± 1.3	1.4 ± 0.5	2.8 ± 0.7	2.8 ± 0.6	5.2 ± 0.9
4	0.9 ± 0.5	3.1 ± 0.8	0.3 ± 0.2	3.8 ± 0.8	0.6 ± 0.3	2.5 ± 0.6	0.00	5.5 ± 0.9
⟨MLs⟩	2.0 ± 0.3	5.8 ± 0.3	3.0 ± 0.4	8.0 ± 0.6	2.5 ± 0.6	2.3 ± 0.6	4.0 ± 0.8	6.3 ± 1.0
$T_{\text{ice}} = 60$ K								
1	8.8 ± 1.0	4.8 ± 0.9	5.3 ± 0.9	11.0 ± 1.3	3.9 ± 0.8	10.0 ± 1.3	3.1 ± 0.7	21.5 ± 1.8
2	2.7 ± 1.2	4.7 ± 0.9	3.2 ± 0.7	20.0 ± 1.8	3.9 ± 0.8	10.8 ± 1.3	4.6 ± 0.8	15.4 ± 1.5
3	1.5 ± 0.9	7.3 ± 1.2	3.0 ± 0.7	23.5 ± 1.9	1.1 ± 0.4	2.4 ± 0.6	4.1 ± 0.8	9.2 ± 1.1
4	0.00	0.8 ± 0.5	0.3 ± 0.2	4.3 ± 0.9	0.00	1.6 ± 0.5	0.00	0.9 ± 0.4
⟨MLs⟩	3.3 ± 0.4	4.4 ± 0.4	2.9 ± 0.3	14.0 ± 0.8	2.2 ± 0.6	6.2 ± 1.0	2.9 ± 0.7	12.0 ± 1.3

**Notes.** All values should be multiplied by  $10^{-3}$ . The notation H<sub>2</sub>O\* and HOD\* denote photoexcitation of H<sub>2</sub>O and HDO in H<sub>2</sub>O ice in which the H atom resulting from photodissociation kicks out a neighbouring H<sub>2</sub>O molecule. The notation D<sub>2</sub>O\* and DOH\* indicate the analogous process in which the resulting D atom kicks out a neighbouring H<sub>2</sub>O molecule.

$$P_{\text{KO}}(\text{HDO}; \text{HDO}^*) = P_{\text{KO}}(\text{H}_2\text{O}; \text{HDO}^*) \quad (\text{C.10})$$

The right hand values of Eqs. C.8 and C.9 can be directly obtained from Tables 5 and A.3.  $P_{\text{KO}}(\text{HDO}; \text{HDO}^*)$  can be computed using the approximation in Eq. C.10 and using Eq. C.7 and Tables 5 and A.3.

Using Eqs. C.8–C.10, Eq. C.5 can be rewritten as

$$P_{\text{HDO}}^{\text{KOdes}} \approx r_{\text{HDO}} \times r_{\text{HDO}} \times P_{\text{KO}}(\text{H}_2\text{O}; \text{HDO}^*) + (1 - r_{\text{HDO}}) \times r_{\text{HDO}} \times P_{\text{KO}}(\text{H}_2\text{O}; \text{H}_2\text{O}^*) \quad (\text{C.11})$$

Most importantly, for 10 and 20 K we have approximately that (see Tables 5 and A.3)

$$P_{\text{KO}}(\text{H}_2\text{O}; \text{HDO}^*) \approx P_{\text{KO}}(\text{H}_2\text{O}; \text{H}_2\text{O}^*). \quad (\text{C.12})$$

Inserting Eq. C.12 in Eq. C.11 yields

$$P_{\text{HDO}}^{\text{KOdes}} \approx r_{\text{HDO}} \times P_{\text{KO}}(\text{H}_2\text{O}; \text{H}_2\text{O}^*) \quad (\text{C.13})$$

and inserting Eq. C.12 in Eq. C.6 yields

$$P_{\text{H}_2\text{O}}^{\text{KOdes}} = (1 - r_{\text{HDO}}) \times P_{\text{KO}}(\text{H}_2\text{O}; \text{H}_2\text{O}^*) \quad (\text{C.14})$$

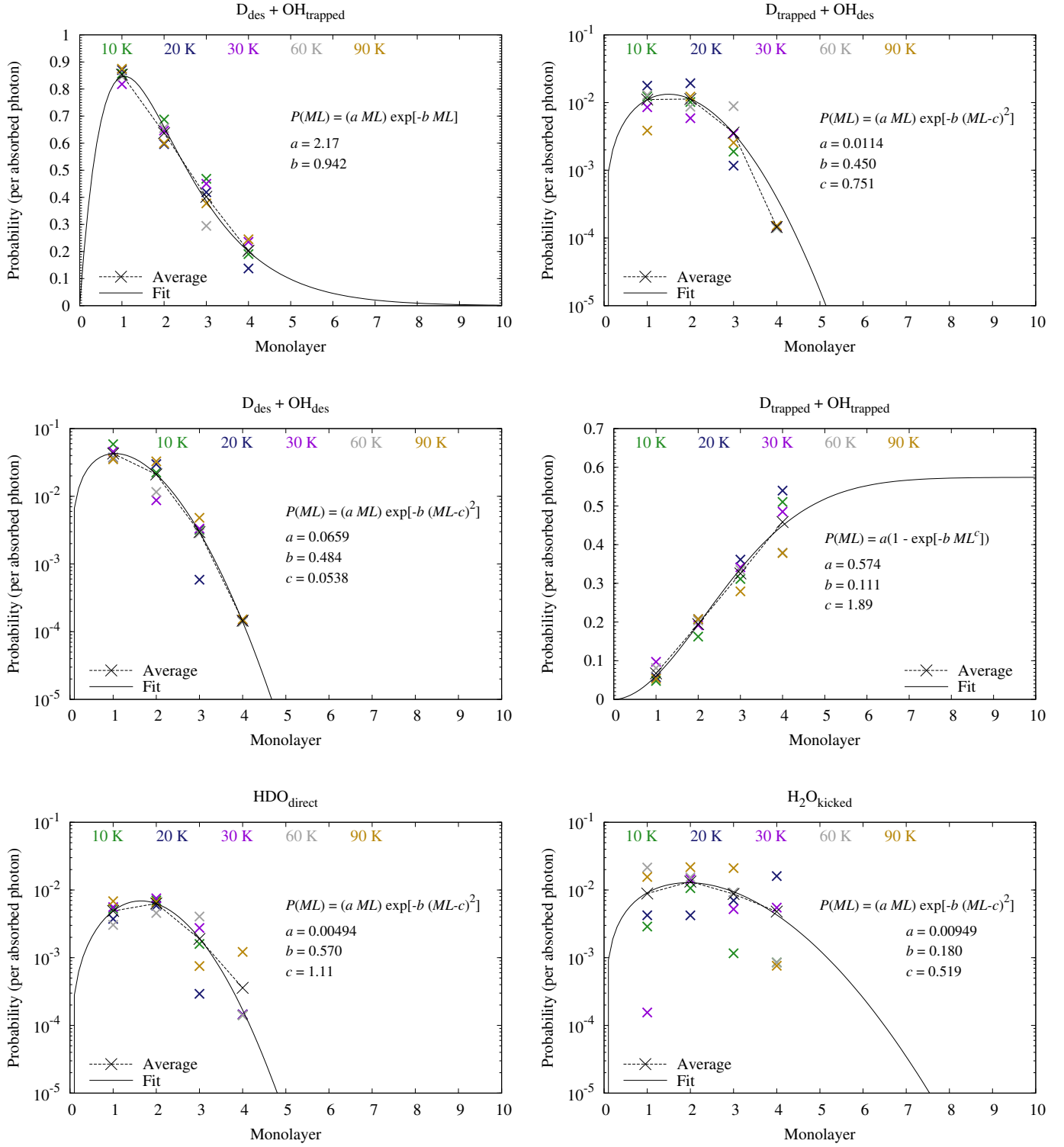
From Eqs. C.13 and C.14, we can derive that

$$P_{\text{HDO}}^{\text{KOdes}} / P_{\text{H}_2\text{O}}^{\text{KOdes}} = r_{\text{HDO}} / r_{\text{H}_2\text{O}} \quad (\text{C.15})$$

meaning that there should be no isotope fractionation due to the indirect kick-out mechanism. Taken together, Eqs. C.4 and C.15 ensure that the ratio of desorbed HDO over desorbed H<sub>2</sub>O in the ice is given by

$$P_{\text{HDO}}^{\text{des}} / P_{\text{H}_2\text{O}}^{\text{des}} = r_{\text{HDO}} / r_{\text{H}_2\text{O}} \quad (\text{C.16})$$

which means that this ratio is simply equal to the ratio of HDO and H<sub>2</sub>O in the ice. Therefore, isotope fractionation does not occur for HDO and H<sub>2</sub>O photodesorption.



**Fig. B.1.** Temperature-specific probabilities, temperature-averaged probabilities, and fitted functions for each outcome as a function monolayer for HDO photodissociation into D + OH.